



Introduction to Seasonal Forecast

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Contents

- What a seasonal forecast is
- Brief introduction to climate dynamics
- Methodology and practices for JMA's seasonal forecast







What a seasonal forecast is

Classification of weather and climate forecasts

According to WMO's Manual on Global Data-processing and Forecasting System,

Class	Target forecast range
Nowcasting	Current and forecasted weather up to 2 hours ahead
Very short-range weather forecasting	Up to 12 hours ahead
Short-range forecasting	From 12 hours to 72 hours ahead
Medium-range weather forecasting	From 72 hours to 240 hours ahead
Extended-range weather forecasting	From 10 days to 30 days ahead, usually averaged and expressed as a departure from climate values for that period
Long-range forecasting	From 30 days up to two years (Monthly, three-month or seasonal outlook of averaged weather parameters)
Climate forecasting	Beyond 2 years (Annual, decadal and beyond, including human-induced climate change projection)



Images of short-range forecasts and long-range forecasts

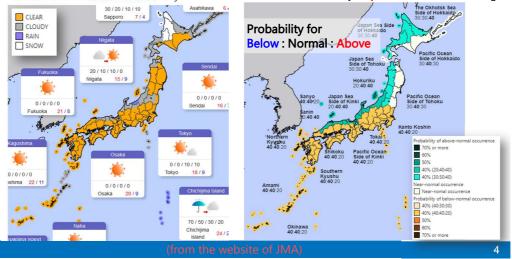
Short-range forecast

States weather parameters (temperatures, precipitation, ...) as they are expected

• Achievable in deterministic way

Seasonal forecast

- States expected deviations from climate values
- Achievable only in **probabilistic** forecasting



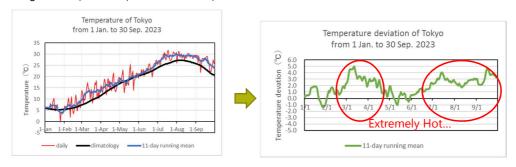


Anomaly is descripted in long-range forecast

Climatological normal is defined as 30-year average for 1991–2020 **Anomaly** is deviation from the Climatology

[Anomaly] = [Actual Value] - [Normal]

- Climate is what we expect, anomaly is what we forecast.
- Anomalies often matter most to industries, societies and economies, because unseasonable weather conditions could bring adverse effects across multiple sectors, including agriculture, tourism, water resource, and so on.

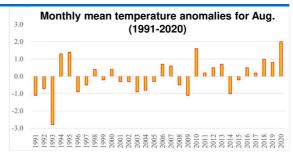


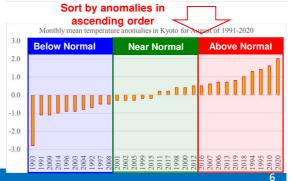


3-category probabilistic forecast

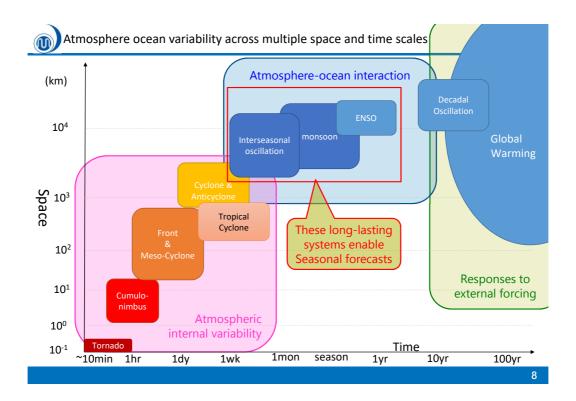
- JMA's seasonal forecasts state probabilities of 3 categories, namely, Above/Near/Below normal
- The 3 categories are derived from historical observations for the 30-year period from 1991 to 2020, by sorting them in ascending order and dividing into 3 categories.
- Seasonal forecasts state probabilities of weather parameter (e.g. temperature) anomalies falling within
 - 1-10th (Below normal; BN)
 - 11-20th (Near normal; NN)
 - 21-30th (Above normal; AN)







Brief introduction to climate dynamics



Anomaly Correlation Coefficient

Anomaly Correlation Coefficient (ACC) is one of measures in the verification of spatial fields (Jolliffe and Stephenson 2003), and is the correlation between anomalies of forecasts and those of verifying values with the reference values, such as climatological values. ACC is defined as follows:

 $ACC = \frac{\sum_{i=1}^{n} w_i(r_i - \overline{f})^2 \sum_{i=1}^{n} w_i(a_i - \overline{a})^2}{\sqrt{\sum_{i=1}^{n} w_i(f_i - \overline{f})^2 \sum_{i=1}^{n} w_i(a_i - \overline{a})^2}}, \quad (-1 \le ACC \le 1),$ where n is the number of samples, and r_i , r, a_i and a are given by the following equations:

 $f_i = F_i - C_i$, $\overline{f} = \left(\sum_{i=1}^n w_i f_i\right) / \sum_{i=1}^n w_i$, $a_i = A_i - C_i$, $\overline{a} = \left(\sum_{i=1}^n w_i a_i\right) / \sum_{i=1}^n w_i$,

where F_i , A_j , and C_i represent forecast, verifying value, and reference value such as climatological value, respectively. Also, f is the mean of f_i , f is the mean of f_i , and f is the mean of f is th of 1. In turn, if the variation pattern is completely reversed, ACC takes the minimum value of -1.

* $w_i = \frac{1}{n}$ (or $\cos \phi_i$, and so on)

Perfect skill Perfectly reversed No ski



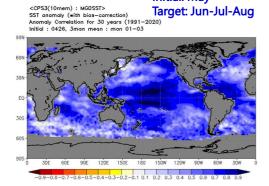
Anomaly Correlation Coefficient of CPS3 products

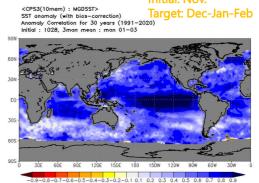
In general,

- Better forecast skill in tropics than in higher latitudes
- Better forecast skill in boreal winter than in summer

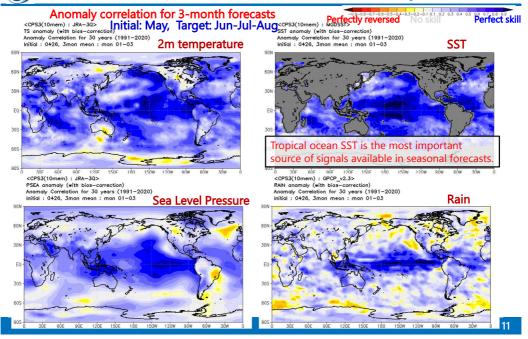
Initial: May

Anomaly correlation for Sea Surface Temperatures for 3-month forecasts Initial: Nov.





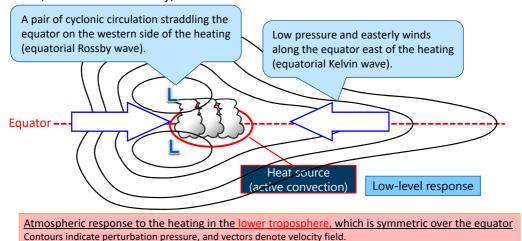
Verification of JMA Seasonal Prediction System





Matsuno-Gill response (in the lower troposphere)

• Gill (1980) found how the tropical atmosphere responds to diabatic heating (i.e. convective activity).



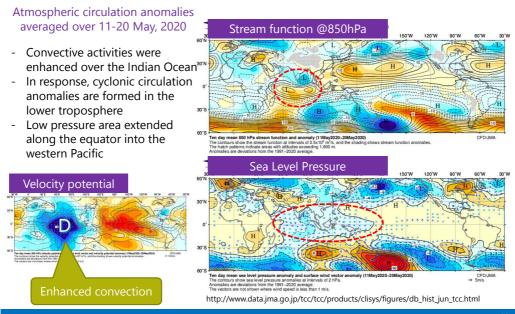
Upper-level response shows the reverse of the low-level response.

Red circle indicates the position of the heating

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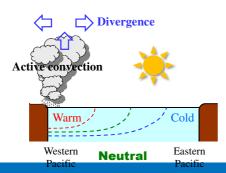
Example of Matsuno-Gill response

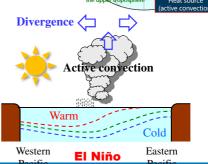




El Nino Southern Oscillation

- During an El Niño event, SST over the central to eastern equatorial Pacific gets warmer significantly.
- The warmer SSTs induce active convection to shift eastward along the equator.
- These convection anomalies give rise to Rossby waves (Matsuno-Gill response!)
- Rossby waves propagate over a large distance and influence the global atmosphere



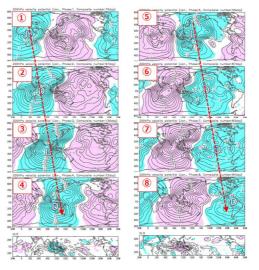


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Madden-Julian Oscillation (MJO)

- MJO is a planetary scale wave consisting of enhanced and suppressed convection extending east-west along the equator
- MJO is the most dominant signal over the tropics on weekly to monthly timescale.
- MJO propagates eastward along the equator, going around the globe in 30 – 60 days.
- In response to MJO, circulation anomalies form and propagate poleward or eastward. This provides key to onemonth forecasts
- MJO is monitored with 200hPa velocity potential (upper-level divergence) field

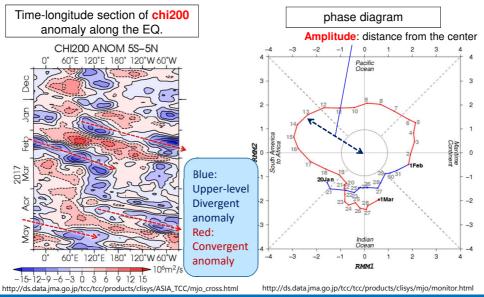


Enhanced phases of convection are denoted as "D" in this figure



How to detect MJO?

Propagation of MJO is visualized through Hovmöller diagram and phase diagram

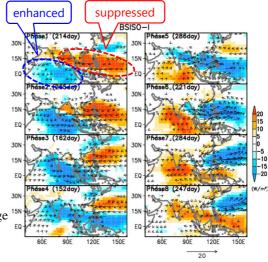


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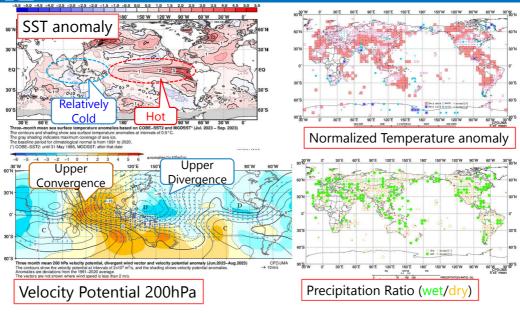
Another "Oscillation" - BSISO

- During summer in northern hemisphere, enhanced or suppressed convection is seen to propagate northward, instead of eastward, over the Indian Ocean and the western Pacific.
- This is called "Boreal Summer Intra-Seasonal Oscillation".
- BSISO can have as much impact on weather conditions as MJO across Asia.
- This is another factor key to extended-range forecasts.



typical time evolution of BSISO





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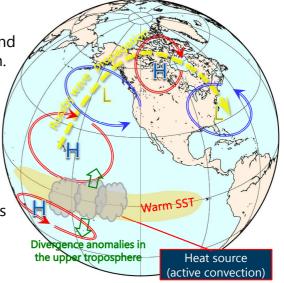
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Matsuno-Gill response in the upper troposphere

 In the upper troposphere, an anti-cyclonic circulation anomaly forms to the north and south of enhanced convection.

 The circulation anomaly propagates poleward as a Rossby wave train.

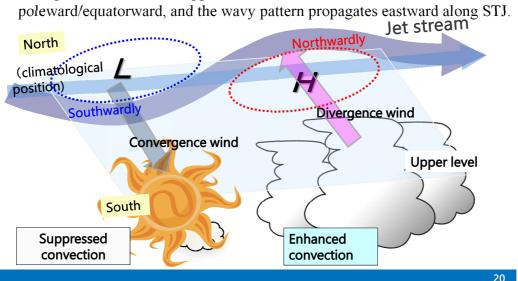
 This sometimes causes anomalous weather conditions in remote areas in subtropics and higher latitudes.





Convection changes jet stream

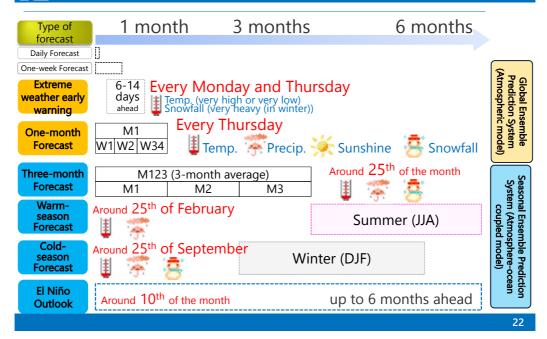
- The subtropical jet stream (STJ) flows poleward of tropics.
- In response to enhanced/suppressed convection, STJ meanders



Methodology for JMA's seasonal forecast



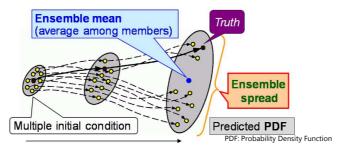
JMA's extended and long-range forecast models





Concept of ensemble prediction

- The atmosphere and ocean is a chaotic system
- Even the tiniest error in an initial condition grows rapidly and errors are unavoidable
- This nature disrupts deterministic numerical prediction beyond about two weeks
- To produce a seasonal forecast, "ensemble prediction" is indispensable.
- Ensemble prediction system (EPS) starts with similar, but slightly different, multiple initial conditions, and produces multiple forecasts.
- With the results from EPS, we can get ensemble mean as the most likely atmospheric conditions in future, along with ensemble spread, an estimation of degree of uncertainty.

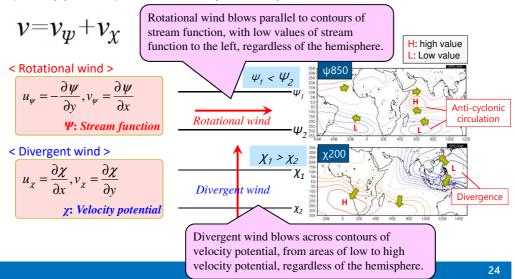


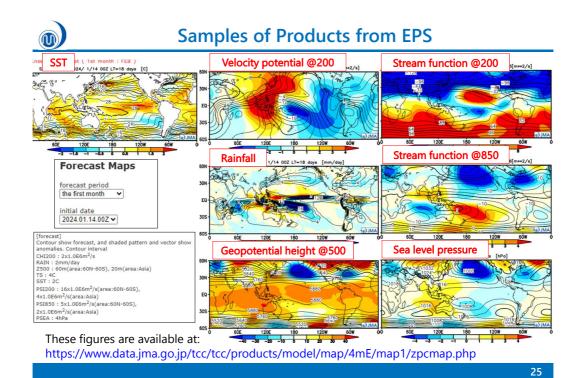
The individual calculation is called "<u>Ensemble member"</u> and the standard deviation among all members is called "<u>Ensemble spread</u>".



Stream function and velocity potential

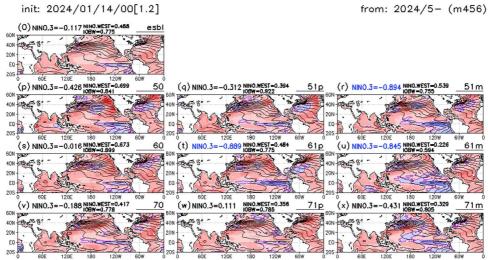
- In discussing seasonal forecast, we often encounter these figures.
- Decomposing wind into a rotational part (stream function) and a divergent part (velocity potential) is useful to analyze atmospheric circulation.







Sample of individual ensemble members



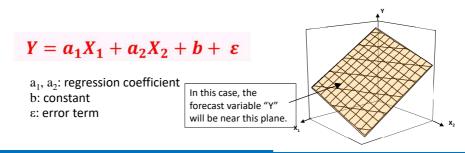
Sea Surface Temperature of the ensemble mean and 9 members (3 days, 3 members per a day) out of 51 (17 days : initialized on 28,Dec \sim 14, Jan.) and its anomaly (shaded) forecasted for May-Jun-Jul 2024 .

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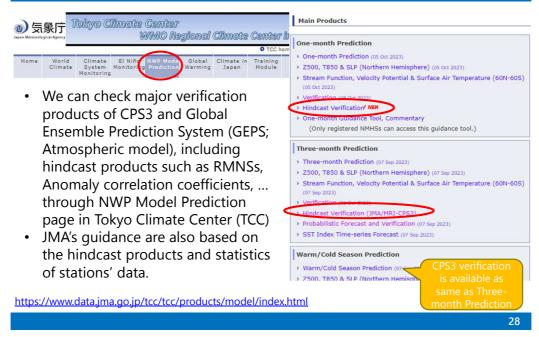
Statistical Post-processing for Seasonal forecast

- Statistical post-processing methods and tools which translate direct model outputs into variables seen in forecasts (e.g. temperature, precipitation, ...) is referred to as "guidance" in JMA.
- Guidance tools exploit statistical relations between past forecasts and observations.
- Typically, these statistical relations are represented in multiple regression equation.

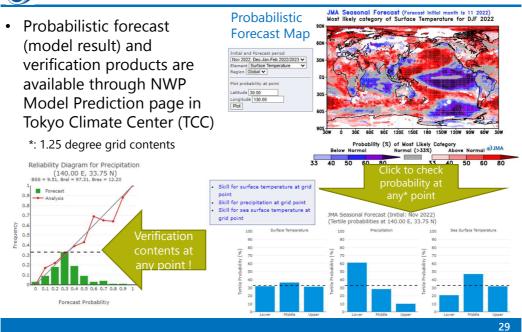




Verification of JMA Seasonal Prediction System



Probabilistic Forecast and Verification contents of TCC



The process to produce seasonal forecast in JMA

